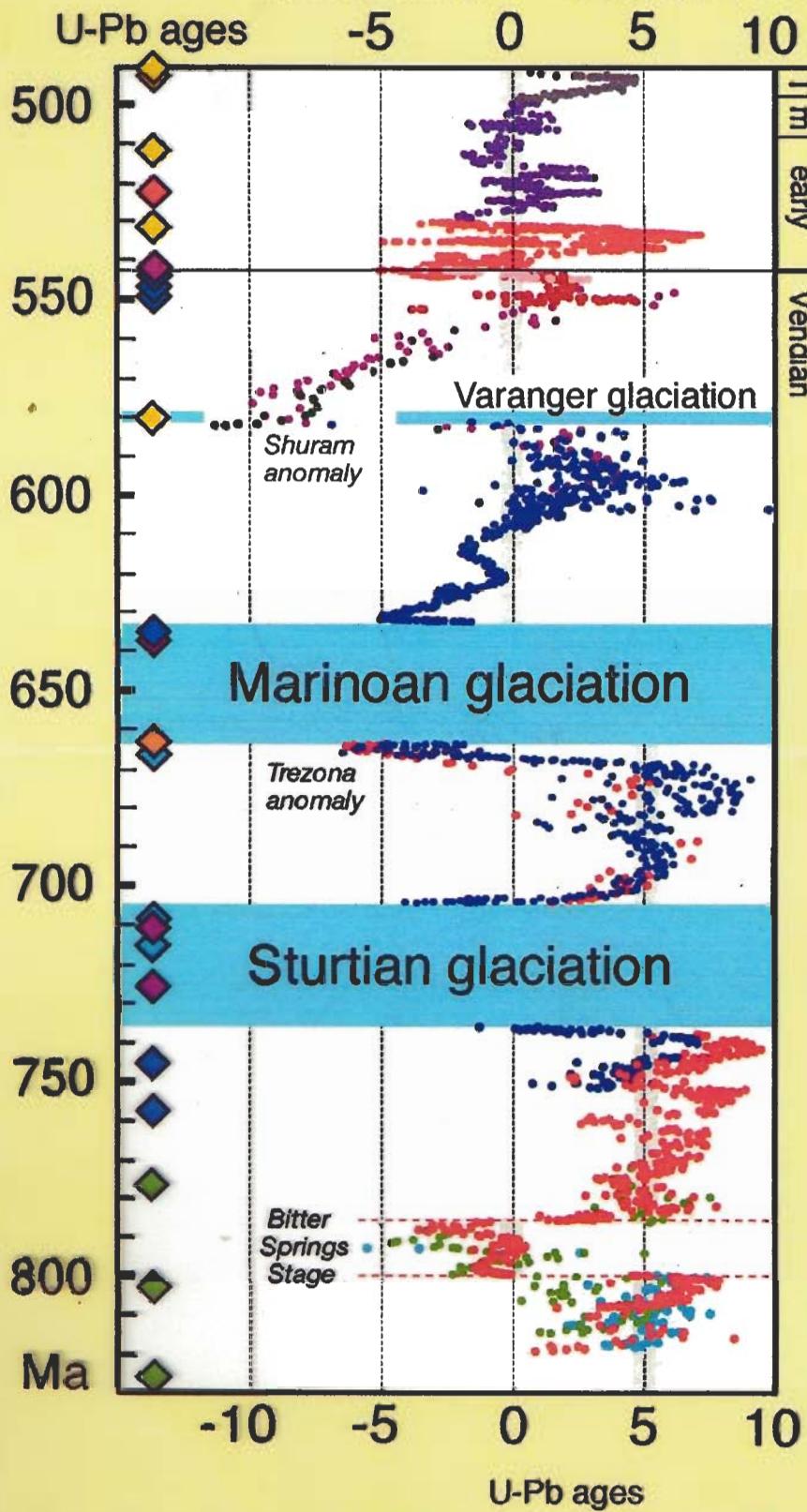


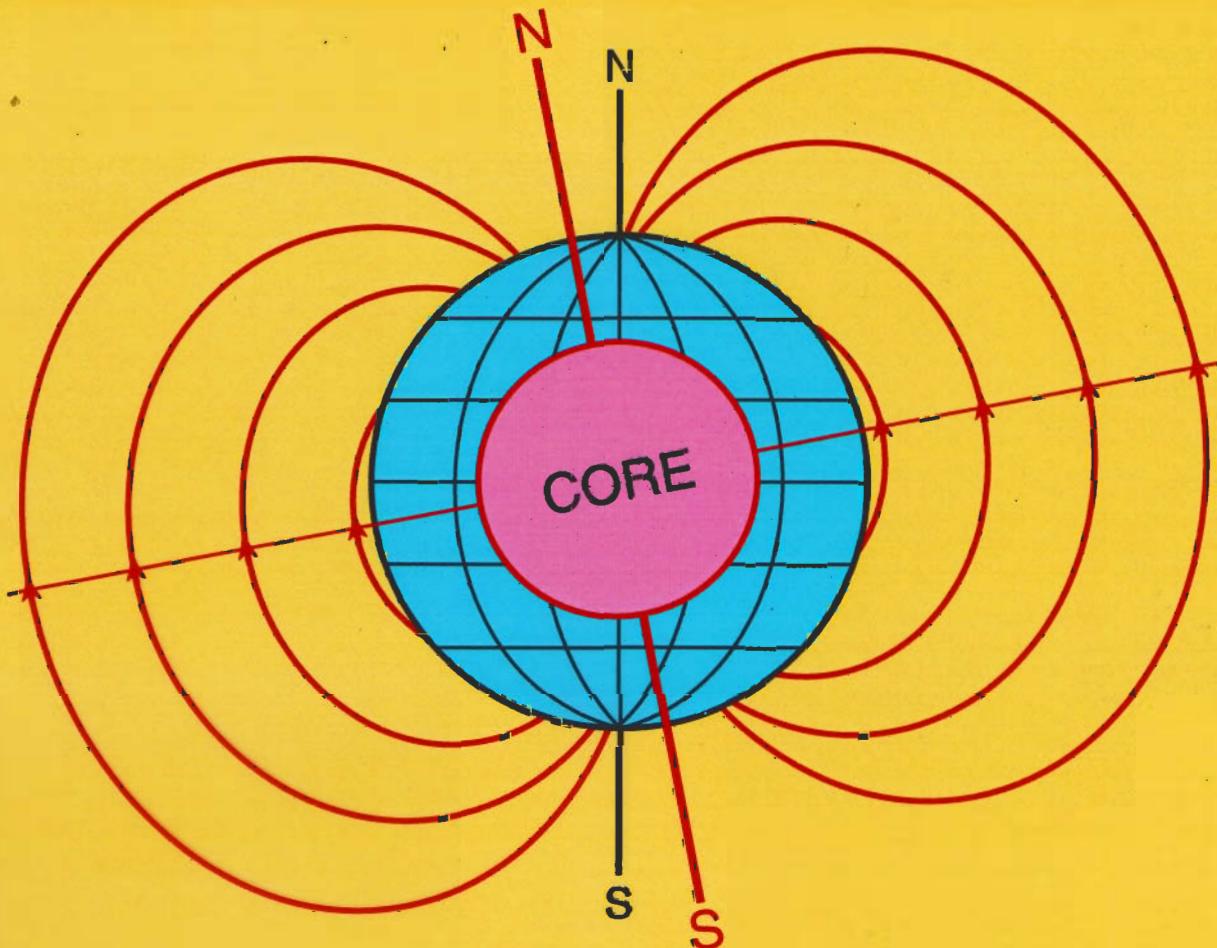
Seawater proxy $\delta^{13}\text{C}_{\text{carb}}$ (VPDB)

after Halverson et al.
(2003)



- Great Basin, USA
Saltzman et al. (2000)
- Yudoma-Olenek, Siberia
Brasier and Sukhov (1998)
- Adoudounian Fm, Morocco
A.C. Maloof (unpubl.)
- Turkut Fm, Siberia
Bartley et al. (1998)
- Nama Gp, Namibia
Saylor et al. (1998)
- Wonoka Fm, Australia
Calver (2000)
- Nafun Group, Oman
Burns and Matter (1993)
- Otavi Group, Namibia
Halverson and Hoffman (2003)
- Hecla Hoek Sgp, Svalbard
Halverson (2003)
- Bitter Springs Fm, Australia
Hill and Walter (2000)
- Shaler Group, Canada
Asmerom et al. (1991)

◆ Australia ◆ Avalonia ◆ China ◆ Morocco ◆ Namibia ◆ Oman ◆ USA



THE EARTH'S DIPOLEAR MAGNETIC FIELD

Radiative heat balance (at equilibrium):

HEAT ABSORBED = HEAT EMITTED

$$R^2 E_s [1 - \alpha] = 4 R^2 [f \sigma T_s^4]$$

R = radius of the Earth

E_s = solar irradiance

α = planetary albedo

f = effective infrared transmission factor (greenhouse effect)

σ = Stefan-Boltzman constant

T_s = surface temperature

PLANETARY ALBEDO: The fraction of incoming radiation that is reflected back to space.

[sea water ~0.1; bare land ~0.3; sea ice ~0.6; fresh snow ~0.9]

ICE-ALBEDO FEEDBACK: For any imposed cooling (or warming), the resulting higher (or) albedo will cause further cooling (or warming). Thus, ice advance is self-stabilizing.

RUNAWAY ICE ALBEDO: If ice lines close to within ~30° of the equator, the ice albedo feedback becomes unstoppable and ice quickly covers the tropics.

see Budyko, M.I., *The effect of solar radiation variations on the climate of the Earth; TELLUS 21: 611-619 (1969).*